



Title: NEON Algorithm Theoretical Basis Document (ATBD): Groundwater Temperature, Elevation and Specific Conductance		Date: 12/06/2019
NEON Doc. #: NEON.DOC.001328	Author: J. Vance	Revision: B

# ALGORITHM THEORETICAL BASIS DOCUMENT (ATBD): GROUNDWATER TEMPERATURE, ELEVATION AND SPECIFIC CONDUCTANCE

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## Change Record

<b>REVISION</b>	<b>DATE</b>	<b>ECO #</b>	<b>DESCRIPTION OF CHANGE</b>
A	08/10/2016	ECO-03872	Initial Release
B	12/06/2019	ECO-05976	Update to sensor elevation algorithm and information

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**1 DESCRIPTION**

Contained in this document are details concerning groundwater temperature, elevation (stage), and specific conductance measurements made in groundwater observation wells at all NEON Aquatic sites. Specifically, the processes necessary to convert “raw” sensor measurements into meaningful scientific units and their associated uncertainties are described. Groundwater temperature, elevation and specific conductance will be continuously monitored by NEON at core and relocatable Aquatic sites using a single sensor (In-Situ, Inc. Aqua Troll 200, AT200).

**1.1 Purpose**

This document details the algorithms used for creating NEON Level 1 data products for groundwater temperature, elevation, and specific conductance from Level 0 data, and ancillary data as defined in this document (such as calibration data) obtained via instrumental measurements made by the Aqua Troll 200 (here after AT200) It includes a detailed discussion of measurement theory and implementation, appropriate theoretical background, data product provenance, quality assurance and control methods

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used, approximations and/or assumptions made, and a detailed exposition of uncertainty resulting in a cumulative reported uncertainty for this product.

## 1.2 Scope

The theoretical background and entire algorithmic process used to derive Level 1 data from Level 0 data for AT200 is described in this document. The sensor employed is a multiparameter probe, the Aqua Troll 200 (AT200), manufactured by In-Situ Inc. (Manufacturer PN: 005610), which measures temperature, pressure and the actual conductivity of the water- in this case groundwater. This document does not provide computational implementation details, except for cases where these stem directly from algorithmic choices explained here.

## 2 RELATED DOCUMENTS, ACRONYMS AND VARIABLE NOMENCLATURE

### 2.1 Applicable Documents

AD[01]	NEON.DOC.000001	NEON OBSERVATORY DESIGN
AD[02]	NEON.DOC.005003	NEON Scientific Data Products Catalog
AD[03]	NEON.DOC.005004	NEON Level 1-3 Data Products Catalog
AD[04]	NEON.DOC.005005	NEON Level 0 Data Products Catalog
AD[05]	NEON.DOC.000782	ATBD QA/QC Data Consistency
AD[06]	NEON.DOC.011081	ATBD QA/QC plausibility tests
AD[07]	NEON.DOC.000783	ATBD De-spiking and time series analyses
AD[08]	NEON.DOC.000746	Evaluating Uncertainty (CVAL)
AD[09]	NEON.DOC.000785	TIS Level 1 Data Products Uncertainty Budget Estimation Plan
AD[10]	NEON.DOC.000751	CVAL Transfer of standard procedure
AD[11]	NEON.DOC.000927	NEON Calibration and Sensor Uncertainty Values
AD[12]	NEON.DOC.001113	Quality Flags and Quality Metrics for TIS Data Products

### 2.2

### 2.3 Reference Documents

RD[01]	NEON.DOC.000008	NEON Acronym List
RD[02]	NEON.DOC.000243	NEON Glossary of Terms

### 2.4

### 2.5 Acronyms

Acronym	Explanation
AIS	Aquatic Instrument System
AT200	Aquat Troll 200
ATBD	Algorithm Theoretical Basis Document
CI	NEON Cyberinfrastructure
CVAL	NEON Calibration, Validation, and Audit Laboratory
DAS	Data Acquisition System
DP	Data Product
FDAS	Field Data Acquisition System
GRAPE	Grouped Remote Analog Peripheral Equipment
Hz	Hertz
L0	Level 0
L1	Level 1
PRT	Platinum resistance thermometer
QA/QC	Quality assurance and quality control

## 2.6 Variable Nomenclature

The symbols used to display the various inputs in the ATBD, e.g., calibration coefficients and uncertainty estimates, were chosen so that the equations can be easily interpreted by the reader. However, the symbols provided will not always reflect NEON's internal notation, which is relevant for CI's use, and or the notation that is used to present variables on NEON's data portal. Therefore a lookup table is provided in order to distinguish what symbols specific variables can be tied to in the following document.

Symbol	Internal Notation		Description
$C_{T0}$	CVALA0		Calibration coefficient for AT200 Temperature sensor
$C_{T1}$	CVALA1		Calibration coefficient for AT200 Temperature sensor
$C_{T3}$	CVALA2		Calibration coefficient for AT200 Temperature sensor
$C_{P0}$	CVALB0		Calibration coefficient for AT200 Pressure sensor
$C_{P1}$	CVALB1		Calibration coefficient for AT200 Pressure sensor
$C_{P2}$	CVALB2		Calibration coefficient for AT200 Pressure sensor
Conductivity has split calibration range:	IF X < 100 $\mu$ S	IF X > 100 $\mu$ S	
$C_{C0}$	CVALM0	CVALH0	Calibration coefficient for AT200 Conductivity sensor
$C_{C1}$	CVALM1	CVALH1	Calibration coefficient for AT200 Conductivity sensor
$C_{C2}$	CVALM2	CVALH2	Calibration coefficient for AT200 Conductivity sensor

$u_{A1,C}$	U_CVALA1	Combined, standard calibration uncertainty of the conductivity measurement by AT200 sensor ( $\mu\text{S}$ )
$u_{A1,T}$	U_CVALA1	Combined, standard calibration uncertainty of the temperature measurement by AT200 sensor ( $^{\circ}\text{C}$ )
$u_{A1,P}$	U_CVALA1	Combined, standard calibration uncertainty of the pressure measurement by AT200 sensor (kPa)
$u_{A3,C}$	U_CVALA3	Combined, relative uncertainty (truth and trueness only) of the conductivity measurement by AT200 sensor (%)
$u_{A3,T}$	U_CVALA3	Combined, relative uncertainty (truth and trueness only) of the temperature measurement by AT200 sensor (%)
$u_{A3,P}$	U_CVALA3	Combined, relative uncertainty (truth and trueness only) of the pressure measurement by AT200 sensor (%)
$E_{Sensor,i}$	Named Location Database: "ELEVATION" + "Z_OFFSET" for each groundwater well location	Elevation of the sensor $i^{\text{th}}$ groundwater well (m – above sea level)
$u_c(E_{Sensor,i})$	Named Location Manager: "ELEVATION_UNCERTAINTY" for each groundwater well location	Combined uncertainty of elevation of the sensor (m)

### 3 DATA PRODUCT DESCRIPTION

#### 3.1 Variables Reported

The groundwater related L1 DPs provided by the algorithms documented in this ATBD are displayed in the accompanying files: gwe\_datapub\_NEONDOC001328.txt, gwc\_datapub\_NEONDOC999999.txt, and gwt\_datapub\_NEONDOC999999.txt.



### 3.2 Input Dependencies

Table 1 details the AT200-related L0 DPs used to produce L1 groundwater DPs in this ATBD.

**Table 3-1: List of AT200-related L0 DPs that are transformed into L1 groundwater DPs in this ATBD.**

Description	Sample Frequency	Units	Data Product Number
Groundwater Pressure	.00333 Hz	kPa	NEON.DOM.SITE.DP0.20015.001.01376.HOR.VER.000
Groundwater Temperature	.00333 Hz	° C	NEON.DOM.SITE.DP0.20015.001.01374.HOR.VER.000
Groundwater Conductance	.00333 Hz	μS/cm	NEON.DOM.SITE.DP0.20015.001.01371.HOR.VER.000
Elevation of the sensor (above sea level, ASL)	NA	m	Geolocation database

### 3.3 Product Instances

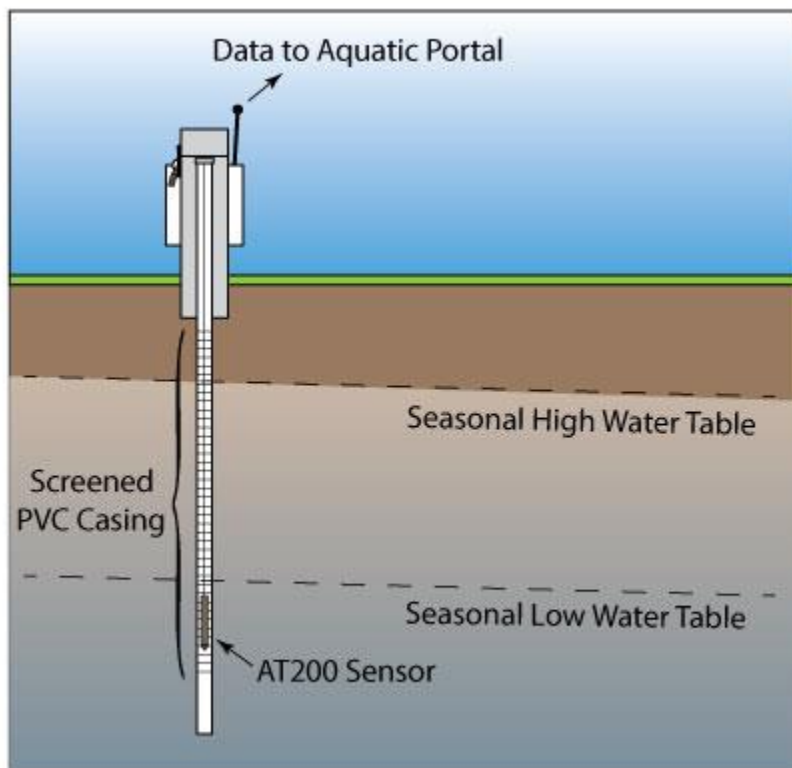
Multiple AT200 sensors will be located at each NEON Aquatic site. AT200 sensors will be located in each of groundwater observation well, hereafter referred to as a well, at NEON Aquatic Sites. Data from each AT200 will be sent to the Aquatic Portal for ingest through wireless data transmission. Special case scenarios will exist in the network where wireless data transmission is not feasible, in which case the data will need to be manually downloaded from each sensor approximately every 2-4 weeks.

### 3.4 Temporal Resolution and Extent

Measurement of temperature, level, and conductivity (TLC) will occur every 5 minutes in each well. The set of AT200's will be programmed to obtain data simultaneously from all wells and will measure and report the specific TLC parameters as a linear average. Each sensor will be internally programmed to collect three individual measurements for each TLC parameter over a 15 second interval centered on the scheduled time of the measurement. The sensor will internally compute the average and will report this value as a single measurement in the data stream from the sensor. Retrieval of the individual measurements is not possible from the sensor so the Level 0 data will be the averaged measurements.

### 3.5 Spatial Resolution and Extent

A single AT200 will be located in each well at each aquatic site. The number of wells will vary between sites due to various reasons, but will range from 6-8 per site, and will be positioned around or on one side of the aquatic feature (stream, river, lake) based on site dependent constraints such as topography or landowner permission. Within each well the AT200 sensor will be positioned near the lower portion of the screened casing of the well, to minimize the potential that the sensor goes above the water table surface in the well. Figure 1 shows a cross-section of a typical well illustrating the positioning of the sensor with respect to the well screen and the water table. Details to sensor function and why this is important are presented in Section 4.1 below.



**Figure 1.** Cross-section of a typical well illustrating the location of the sensor compared to the well screen and the water table, including seasonal variations.

## 4 SCIENTIFIC CONTEXT

Groundwater level, temperature, and specific conductance are fundamental parameters to monitor as a proxy for determining groundwater movement and quality. Groundwater is the water contained in the soils beneath the surface of the earth. Typically this water has fallen on the surface of the earth in some form of precipitation (rain/snow), has percolated into the soils or rocks, and has seeped downward into the earth until it hit a relatively impermeable layer of soil or rock. This layer creates a lower bound for

further infiltrating water to pool up on creating a layer of saturation in the soils and rocks beneath the earth’s surface. Due to gradients (e.g. topographical, pressure) groundwater moves, albeit very slowly, through the soils beneath our feet. The zone of saturated media is bounded at the top by air and the interface between the fully saturated media and the unsaturated media is termed the water table. Technically it’s defined as the point in the vertical soil profile where the hydrostatic pressure of water is zero, since just above this point there a negative hydrostatic pressure due to capillary action, and below this point the depth below water table surface increases and thus the hydrostatic pressure increases.

Water table elevation is the key parameter used in tracking the rate and/or volume of groundwater movement through the subsurface. This is done by comparing the water table elevation between several wells; and relies on accurately knowing the spatial orientation of an array of groundwater wells and sensor positioning in each well. Water flows in the direction of decreasing hydrostatic gradient and as such, knowing the water table elevations in each well and the distance between wells allows for mapping and calculating the seasonal rates and variations in groundwater flow paths, volumes, and movement rates. Knowing groundwater flow rate is critical in the ability to calculate ecosystem processes such as estimating water budgets, calculating nutrient fluxes, or examining surface water groundwater interactions.

Groundwater temperature plays an important role in regulating surface water temperature in streams, which is critical to supporting aquatic life as many aquatic organisms can only survive in a limited temperature range. Groundwater temperature fluctuates slightly between day and night (diurnal cycling) and more so during seasonal changes but below about 4 meters tends to stay relatively consistent throughout the year.

Groundwater specific conductance is a measure of the concentration of ionic solutes dissolved in the groundwater. This measurement is used as a proxy for water quality since provides a direct measurement of the presence of chemicals in the water and it’s a simple measurement to make. Since pure water is electrically non-conductive, the lower the concentration of ionic chemical species in the water the lower the magnitude of the measurement will be. The following section details how the sensor measures conductivity.

#### 4.1 Theory of Measurement

The AT200 contains three sensors, one for each parameter - level, temperature, and conductance. The AT200 converts the analog measurements to digital signals internally and is the only output option from sensor. Obtaining raw analog signals are not possible from this sensor. The following paragraphs in this section detail how the sensor makes each measurement.

##### 4.1.1 Groundwater Elevation

Groundwater elevation is determined by first knowing the pressure of the water above the sensor measurement point. This is done through the use for a pressure transducer in the AT200, which determines pressure by measuring the slight voltage change that occurs when a variable resistor is

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compressed or deflected due to the hydrostatic pressure of the water. The sensor head is a diaphragm-like impermeable membrane, which is in direct contact with the water on one side and vented to the atmosphere on the other side via a small tube that is contained in power and communication cable. The imbalance between the pressure of the water and the air pressure on the two sides of the membrane cause it to be deflected towards the air side. The membrane is part of an electrical circuit and is made of a variable resistor material, which when deflected by the water pressure changes the resistance of the circuit and causes a drop in electrical voltage in the circuit. Through calibrations these changes in electrical voltage can be equated directly to changes in hydrostatic pressure, when the AT200 is spatially fixed. Since changes in atmospheric pressure will directly result in changes in the hydrostatic pressure at a given point in the water column, venting the sensor to the atmospheric pressure changes eliminates the need to compensate for the pressure of the atmosphere in the measurement stream.

#### 4.1.2 Groundwater Temperature

Temperature is derived from the AT200, which uses a temperature dependent electrically resistive material. The sensor applies a fixed current within the circuit and monitors changes in voltage which are directly induced by the temperature dependence of the resistor. Empirically the voltage changes are correlated to water temperature values internally by the sensor prior to data output.

#### 4.1.3 Groundwater Specific Conductance

Specific conductance is the waters ability to conduct electricity which was detailed above as a proxy for water quality. The higher the concentration of dissolved ions is in the water, the less resistive the water is. The measurement is made by the sensor through a set of 4 equally spaced electrodes near the sensor tip which are in contact with the water. The sensor uses the electrodes to create an electrical circuit in the water and applies a fixed current to the circuit. The electrical conductance of the water is then determined by monitoring the voltage in the circuit after it passes through the water. Electrical conductivity of water is a function of the spacing between the electrodes and the temperature of the system. The sensor has accounted for these parameters by having a spatially fixed electrode placement and by monitoring the water temperature. Actual conductivity is the rawest form of the measurement that the sensor can send to the NEON DAS (GRAPE) and is the conductivity measurement that accounts for electrode spacing but does not correct for temperature effects. Section 4.2 below defines the algorithm to convert actual conductivity (actual measurement) to specific conductivity (temperature corrected).

### 4.2 Theory of Algorithm

The following sections detail the theory of the algorithm for converting water pressure to water table surface elevation (Section 4.2.1) and to normalize conductivity measurements for temperature effects.

#### 4.2.1 Groundwater Surface Elevation Algorithm

Calibrated groundwater pressure will be used to calculate the groundwater elevation. Groundwater pressure will be determined by applying the calibration coefficients, supplied by CVAL, to the “raw” sensor output as follows:

$$P_{GW,i} = C_{P2} * P_i^2 + C_{P1} * P_i + C_{P0} \quad \text{Equation 1}$$

Where:

- $P_{GW,i}$  = Individual (1/300 Hz) groundwater pressure (kPa)
- $C_{P2}$  = Calibration coefficient provided by CVAL ((kPa)<sup>-1</sup>)
- $C_{P1}$  = Calibration coefficient provided by CVAL (unitless)
- $C_{P0}$  = Calibration coefficient provided by CVAL (kPa)
- $P_i$  = Individual (1/300 Hz) pressure output from sensor (kPa)

The elevation of the water table surface is the key point to measuring the hydrostatic pressure in the well and is calculated by knowing the depth from the ground surface at the well to the sensor in the well. The water pressure can be converted to a length term by

$$E_{GW,i} = E_{sensor,i} + 1000 \times (P_{GW,i} / (\rho_{water} \times g)) \quad \text{Equation 2}$$

Where:

- $E_{GW,i}$  = Individual (1/300 Hz) elevation of the water table (m-ASL)
- $E_{sensor,i}$  = Elevation of the sensor within the  $i^{th}$  well (m-ASL) = ELEVATION + Z\_OFFSET in the Named Location Database
- $P_{GW,i}$  = Individual (1/300 Hz) hydrostatic pressure (from sensor) (kPa)
- $\rho_{water}$  = Density of water, 999 (kg/m<sup>3</sup>)
- $g$  = Acceleration due to gravity, 9.81 (m/s<sup>2</sup>)

During the construction of each aquatic site, groundwater wells will be surveyed and the ground surface elevation ( $GS_{well,i}$ ) at each well will be determined. The surveyed location information will be stored in the Named Location Database as an ELEVATION of the survey point. In addition, the depth of the sensor below ground surface will be measured and stored in the Named Location Database as a Z\_OFFSET. These values of the sensor depth are combined to provide the elevation of the sensor ( $E_{sensor,i}$ ). The locations as well as the longitude and latitude for each groundwater well will be stored in the Named Location Database for each NEON AQU site. Based on the design of the wells and AT200 mounting hardware, removal of the sensor should not interfere with the ability to replace the sensor in the same location in the well.

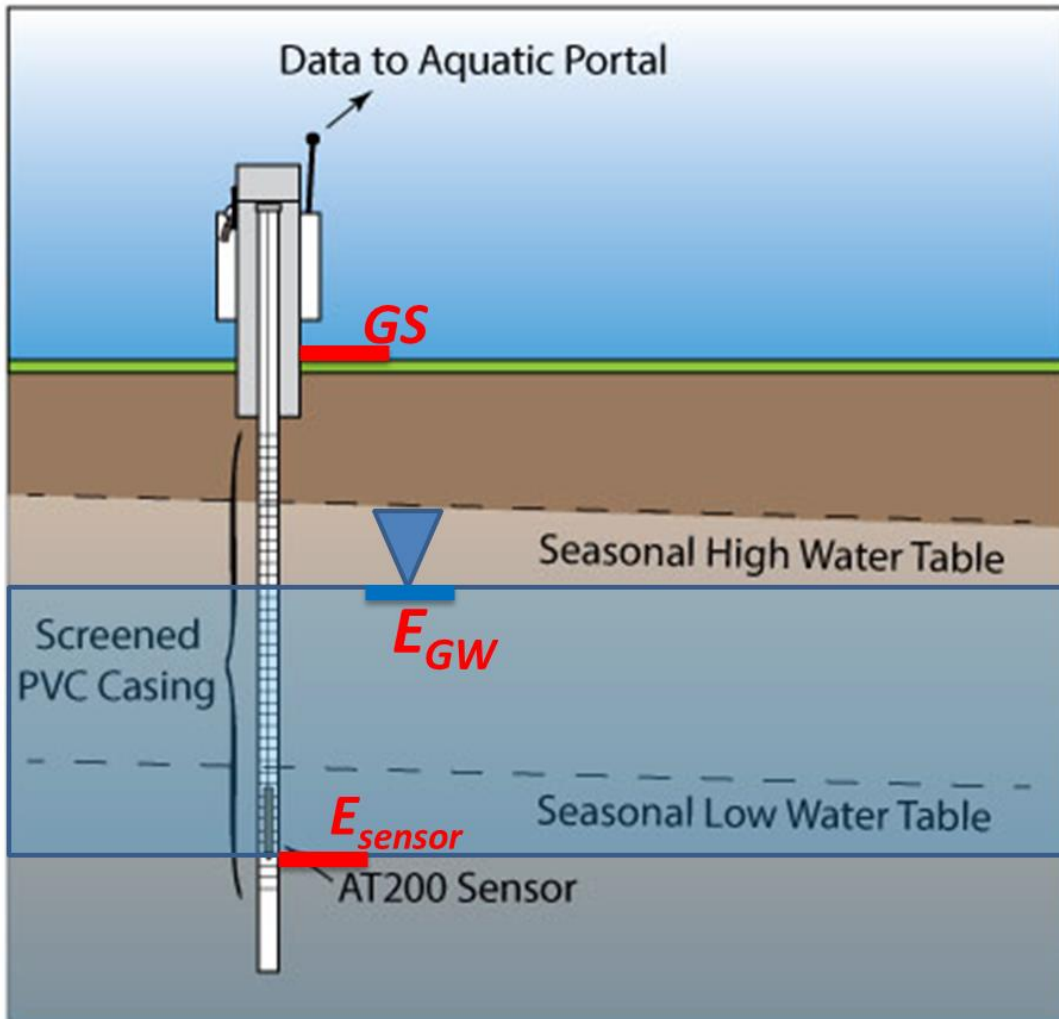


Figure 2. picture showing the reference points for the elevation of the ground surface ( $GS$ ), elevation of the sensor ( $E_{sensor}$ ) and the groundwater elevation ( $E_{GW}$ ).

Groundwater elevation is reported as the instantaneous 5-minute measurement and a 30-minute average. The instantaneous groundwater surface elevation ( $E_{GW,i}$ ) shall be calculated to create the 30-minute average as:

$$\overline{E}_{30} = \frac{1}{n} \sum_{i=1}^n E_{GW,i} \quad \text{Equation 3}$$

Where for each 30-minute averaging,  $n$  is the number of measurements in the averaging period and  $E_{GW,i}$  is the groundwater elevation calculated from the 1/300 Hz pressure measurement according to equations 1-2 above during the 30-minute averaging period. For a 30-minute average,  $n = 6$  if all points are included [0,6).

#### 4.2.2 Groundwater Temperature Algorithm

Groundwater temperature will be reported as the instantaneous 5-minute measurement and a 30-minute average. The instantaneous (0.003 Hz) temperature will be determined accordingly to create additional L1 DPs:

$$T_{GW,i} = C_{T2} * T_i^2 + C_{T1} * T_i + C_{T0} \quad \text{Equation 4}$$

Where:

- $T_{GW,i}$  = Individual (1/300 Hz) groundwater temperature (°C)
- $C_{T2}$  = Calibration coefficient provided by CVAL ((°C)<sup>-1</sup>)
- $C_{T1}$  = Calibration coefficient provided by CVAL (unitless)
- $C_{T0}$  = Calibration coefficient provided by CVAL (°C)
- $T_i$  = Individual (1/300 Hz) temperature output from sensor (°C)

The instantaneous temperature measurements will be used to calculate the 30-minute average according to:

$$\overline{T}_{30} = \frac{1}{n} \sum_{i=x}^n T_{GW,i} \quad \text{Equation 5}$$

Where for each 30-minute averaging,  $n$  is the number of measurements in the averaging period and  $T_{GW,i}$  is the groundwater temperature calculated from the 1/300 Hz temperature measurement according to equations 4 above during the 30-minute averaging period. For a 30-minute average,  $n = 6$  if all points are included [0,6).

#### 4.2.3 Groundwater Specific Conductivity Algorithm

Temperature directly affects the ability of electricity to flow through water. Actual conductivity is the term given to an electrical conductivity measurement of water that is not normalized for temperature effects; whereas specific conductance is the value at a water temperature normalized to 25°C. Normalization of the conductivity measurement allows for comparison of the metric without temperature bias and is calculated as (Wilde et al, various dates).

Actual groundwater conductivity will be determined by applying the calibration coefficients, supplied by CVAL, to the “raw” sensor output as follows:

$$C_{GW,i} = C_{C2} * C_i^2 + C_{C1} * C_i + C_{C0} \quad \text{Equation 6}$$

Specific conductivity will then be calculated from actual conductivity and temperature using the calibrated 5-minute values according to Eq.7.

$$SpC_{GW,i} = \frac{C_{GW,i}}{1+0.0191(T_{GW,i}-25)} \quad \text{Equation 7}$$

where:

$SpC_{GW,i}$  = Individual Conductivity at 25°C, known as Specific Conductance ( $\mu S$ )  
 $C_{GW,i}$  = Individual Conductivity as measured at the ambient temperature ( $\mu S$ )  
 $T_{GW,i}$  = Individual Groundwater (ambient) temperature ( $^{\circ}C$ )

If the synchronous temperature measurements are not available, then specific conductance shall not be calculated and missing data will be flagged.

After conductivity is temperature corrected to produce specific conductance ( $SpC_{GW}$ ), 30-minute average of specific conductance ( $\overline{SpC_{30}}$ ) will be determined accordingly to create L1 DPs:

$$\overline{SpC_{30}} = \frac{1}{n} \sum_{i=x}^n SpC_{GW,i} \quad \text{Equation 8}$$

Where for each 30-minute averaging,  $n$  is the number of measurements in the averaging period and  $SpC_{GW,i}$  is the groundwater specific conductivity calculated from the 1/300 Hz conductivity measurement according to equation 6 above during the 30-minute averaging period. For a 30-minute average,  $n = 6$  if all points are included [0,6).

## 5 ALGORITHM IMPLEMENTATION

Data flow for signal processing of L1 DPs will be treated in the following order.

1. Calibration coefficients will be applied to instantaneous 5-minute groundwater pressure value ( $P_i$ ) according to Equation 1.
2. Calibration coefficients will be applied to instantaneous 5-minute groundwater temperature ( $T_i$ ) according to Equation 3.
3. Calibration coefficients will be applied to instantaneous 5-minute groundwater conductivity ( $C_i$ ) according to Equation 6.
4. Groundwater pressure ( $P_{GW}$ ) will be converted to groundwater Surface Elevation ( $E_{GW}$ ) according to Equation 2.
5. Groundwater conductivity ( $C_{GW}$ ) will be converted to Specific Conductivity ( $SpC_{GW}$ ) according to Equation 7\*.
6. QA/QC Plausibility tests will be applied to the AT200 pressure, temperature and conductivity data streams in accordance with AD[06]. The details are provided below.
7. Signal de-spiking will be applied to the AT200 pressure, temperature and conductivity data stream in accordance with AD[07].
8. 30-minute averages will be calculated for Groundwater Temperature ( $\overline{T_{30}}$ ), Groundwater Surface Elevation ( $\overline{E_{30}}$ ) and Specific Conductivity ( $\overline{SpC_{30}}$ ) according to Equations 3, 5, and 8 respectively.
9. Descriptive statistics, i.e. minimum, maximum, and variance, will be determined for thirty-minute averages.



- Quality metrics, quality flags, and the final quality flag will be produced for thirty-minute average according to AD[16].

\* If the synchronous temperature measurements are not available, then specific conductance shall not be calculated and missing data will be flagged.

**QA/QC Procedure:**

- Plausibility Tests** – All plausibility tests will be determined for groundwater pressure, temperature and conductivity (AD[06]). Test parameters will be provided by AQU and maintained in the CI data store. All plausibility tests will be applied to the sensor’s L0 DP and an associated quality flags (QFs) will be generated for each test.
- Signal De-spiking and Time Series Analysis** – The time series de-spiking routine will be run according to AD[07]. Test parameters will be specified by AQU and maintained in the CI data store. Quality flags resulting from the de-spiking analysis will be applied according to AD[07].
- Placeholder for Consistency Analysis** (see section 7 for future implementation).
- Quality Flags (QFs) and Quality Metrics (QMs) AD[16]** – If a datum fails one of the following tests it will not be used to create a L1 DP: **range, persistence, step, null and gap**.  $\alpha$  and  $\beta$  QFs and QMs will be determined using the flags in Table 3. In addition, L1 DPs will have a QA/QC report and quality metrics associated with each flag listed in Table 3 as well as a final quality flag, as detailed in AD[16]. Ancillary information needed for the algorithm and other information maintained in the CI data store is shown in Table 4.
  - Specific Conductivity** – The calculation of specific conductivity from actual conductivity and temperature shall be done before the QA/QC procedure. If the input temperature value is null or erroneous, the corresponding specific conductivity value should also likely fail the QA/QC tests and produce a final QF flag of 1 and therefore not be used in the development of the L1 DP.

**Table 5-1: Flags associated with AT200 measurements.**

<b>Tests</b>
Range
Persistence
Step
Null

Tests
Gap
Signal Despiking
Alpha
Beta
Final quality flag

**Table 5-2: Information maintained in the CI data store for with AT200**

Tests/Values	CI Data Store Contents
Range	Minimum and maximum values
Persistence	Window size, threshold values and maximum time length
Step	Threshold values
Null	Test limit
Gap	Test limit
Signal Despiking	Time segments and threshold values
Calibration	CVAL sensor specific calibration coefficients
Uncertainty	AD[14]
Final Quality Flag	AD[16]

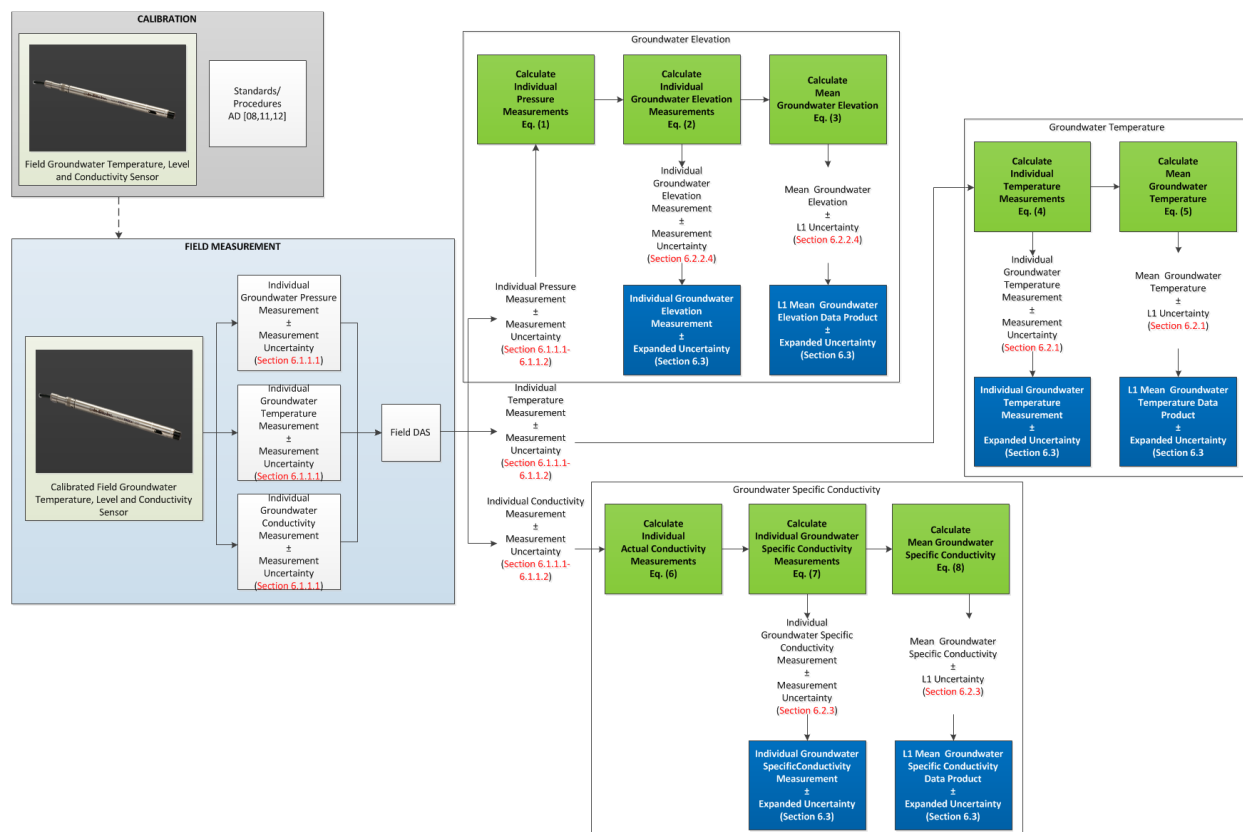
## 6 UNCERTAINTY

Uncertainty of measurement is inevitable; therefore, measurements should be accompanied by a statement of their uncertainty for completeness (JCGM 2008; Taylor 1997). To do so, it is imperative to identify all sources of measurement uncertainty related to the quantity being measured. Quantifying the uncertainty of AIS measurements will provide a measure of the reliability and applicability of individual measurements and AIS data products. This portion of the document serves to identify, evaluate, and quantify sources of uncertainty relating to individual, calibrated groundwater

measurements as well as L1 mean groundwater DPs. It is a reflection of the information described in AD[11], and is explicitly described for the In-Situ AT200 in the following sections.

## 6.1 Uncertainty of Groundwater Measurements

Uncertainty of the AT200 assembly is discussed in this section. The section is broken down into two topics. The first informs the sources of *measurement* uncertainty, i.e., those associated with *individual pressure, temperature and conductivity measurements*. The second details uncertainties associated with temporally averaged data products. A diagram detailing the data flow and known sources of uncertainty are displayed in Figure 3.



**Figure 3.** Displays the data flow and associated uncertainties of individual measurements of groundwater temperature, level (from pressure) and conductivity and associated L1 DPs.

### 6.1.1 Measurement Uncertainty

The following subsections present the uncertainties associated with *individual observations*. It is important to note that the uncertainties presented in the following subsections are *measurement uncertainties*, that is, they reflect the uncertainty of an *individual* measurement. These uncertainties

should not be confused with those presented in Section 6.1.2. We urge the reader to refer to AD[11] for further details concerning the discrepancies between quantification of measurement uncertainties and L1 uncertainties.

NEON calculates measurement uncertainties according to recommendations of the Joint Committee for Guides in Metrology (JCGM) 2008. In essence, if a measurand  $y$  is a function of  $n$  input quantities

$x_i$  ( $i = 1, \dots, n$ ), *i.e.*,  $y = f(x_1, x_2, \dots, x_n)$ , the combined measurement uncertainty of  $y$ , assuming the inputs are independent, can be calculated as follows:

$$u_c(y) = \left( \sum_{i=1}^N \left( \frac{\partial f}{\partial x_i} \right)^2 u^2(x_i) \right)^{\frac{1}{2}} \quad (9)$$

where

$\frac{\partial f}{\partial x_i}$  =partial derivative of  $y$  with respect to  $x_i$

$u(x_i)$  =combined standard uncertainty of  $x_i$ .

Thus, the uncertainty of the measurand can be found by summing the input uncertainties in quadrature. For groundwater pressure, temperature and conductivity measurements, the sources of uncertainty are discussed below.

#### 6.1.1.1 DAS

The In-Situ AT200 sensor has an internal Analog to Digital (A/D) converter and outputs data in digital form. Therefore, no data conversions occur within the DAS, and uncertainty introduced by the DAS can be considered negligible.

#### 6.1.1.2 Calibration

Uncertainties associated with the calibration process of the AT200 for groundwater pressure, temperature and conductivity measurements will be provided by CVAL as individual standard combined uncertainty values. These uncertainties  $\{u_{A1}\}$  (see Section 2.4) represent i) the repeatability and reproducibility of the sensor and the lab DAS and ii) uncertainty of the calibration procedures and coefficients including uncertainty in the standard (truth). Both are constant values that will be provided by CVAL, stored in the CI data store, and applied to all *individual pressure, temperature and conductivity measurements* (that is, the uncertainty values do not vary with any specific sensor, DAS component, etc.). A detailed summary of the calibration procedures and corresponding uncertainty estimates can be found in AD[11,12].

### 6.1.1.3 Groundwater Temperature

There is no additional uncertainty beyond the measurement and calibration uncertainties that needs to be considered.

### 6.1.1.4 Groundwater Surface Elevation

This Section details measurement uncertainty relating to individual *Groundwater Surface Elevation* measurements. Figure 4 shows the variables necessary to calculate the sensor elevation ( $E_{Sensor}$ ). These will be measured by SI&V during sensor installations at each groundwater well.

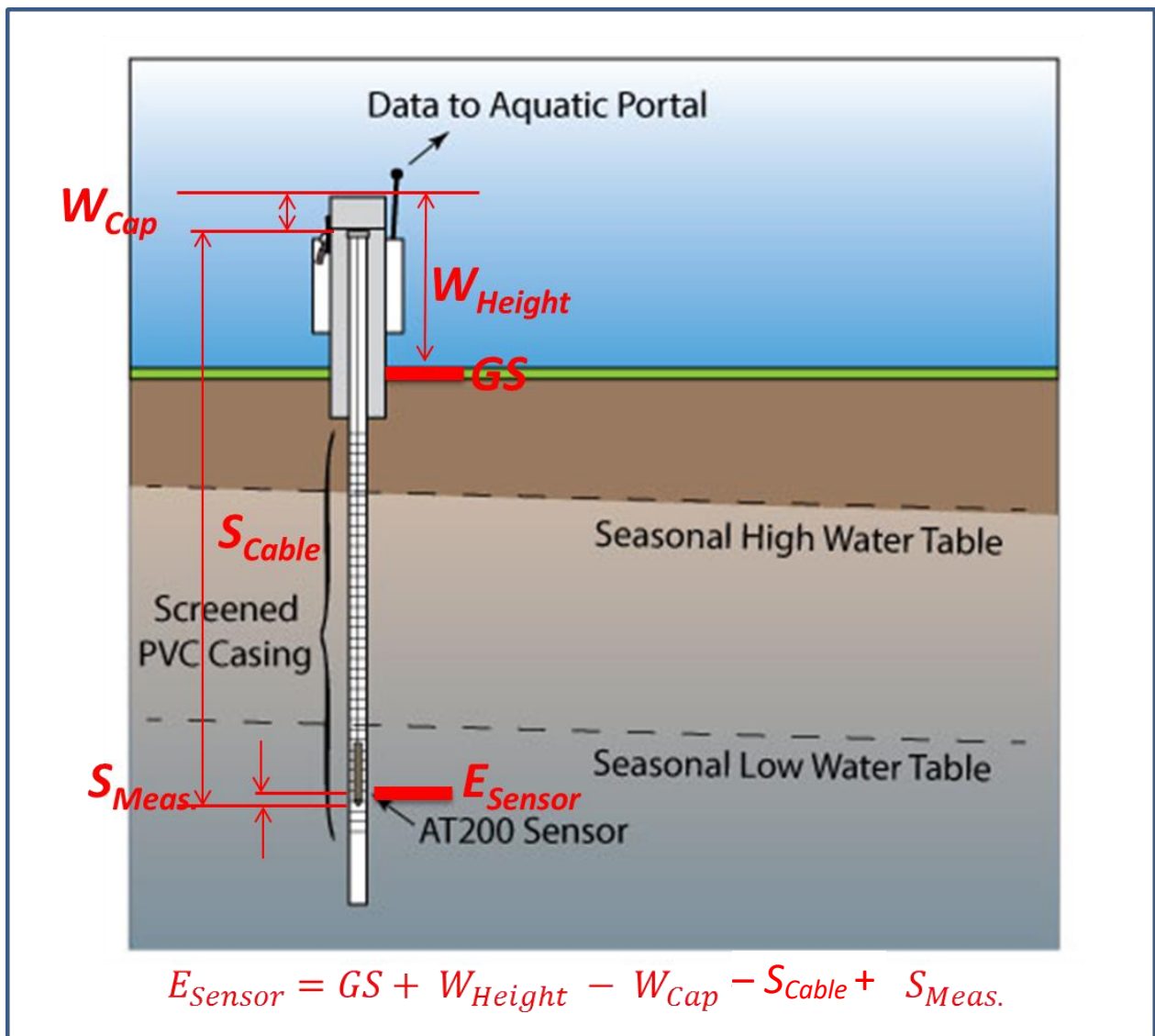


Figure 4. Illustration of the points measured by SI&V during installation and the calculation of the sensor elevation.

#### 6.1.1.4.1 Ground Surface Elevation

Spatial error is a principle source of uncertainty in the calculation of the water table elevation. The precise location and elevation of the groundwater well casing as well as the ground surface elevation at each well will be determined after installation by surveying the site. Spatial error will combine the uncertainty of the permanent survey marker at a site and the error associated with surveying the ground elevation at individual wells. These errors will be quantified for both the permanent benchmark and the ground surface elevation at the well location according to the survey process.

The combined uncertainty of the ground surface elevation for all wells will be calculated by SI&V according to Equation 10 based on the technique used across all sites. This is shown to be informative only. CI will not need to utilize this constant (See next section).

$$u_c(GS) = \left[ u(M)^2 + u(GS_n)^2 \right]^{\frac{1}{2}} \quad \text{Equation 10}$$

Where

$u_c(GS)$  = Combined uncertainty of the ground surface elevation at individual (n) well casings (m)

$u(M)$  = Uncertainty of the measurement of the site marker, reported (m)

$u(GS)$  = Uncertainty of the measurement of the elevation of the ground at the base of the well casing (m)

These values are given as the accuracy of the measurements during surveying and will be passed from the surveyor to SI&V and determined as a constant value for all wells.

#### 6.1.1.4.2 Sensor Depth

Measurement of the depth of the sensor from the ground surface relies upon accurately measuring the distance from the top of the groundwater well casing the ground surface ( $WH$ ) and the distance from the well casing top to the sensor measurement point ( $MD$ ). Uncertainty of these two measurements are additive and the total uncertainty of sensor depth placement will combine the uncertainty of both manual measurements which will be quantified in the field and passed to CI in a metadata file.

$$u_c(E_{sensor,i}) = \left[ u_c(GS)^2 + u(W_{Height})^2 + u(W_{Cap})^2 + u(S_{Cable})^2 + u(S_{Meas.})^2 \right]^{\frac{1}{2}} \quad \text{Equation 11}$$

Where

$u_c(E_{sensor})$  = combined uncertainty of the sensor elevation (m-ASL)

$u(W_{Height})$  = uncertainty of the measurement of the height of the well casing with respect to the ground surface ( $GS$ ) (m)

- $u(W_{Cap})$  = uncertainty of measurement of the distance from the inlet of the well inside the casing (PVC pip) to the top of the well casing
- $u(S_{Cable})$  = uncertainty of measurement of the length of the cable from the top of the inlet of the well inside the casing to the bottom of the sensor body
- $u(S_{Meas.})$  = uncertainty of the measurement of the distance from the point of measurement to the bottom of the sensor body

These values are given as the accuracy of the measurements during surveying and will be provided by SI&V after the time of installation at a site to AQU. The uncertainty of sensor evaluation will be determined from replicate installs and given as constant value for all wells. **The final location of this uncertainty value is yet to be determined.**

#### 6.1.1.4.3 Density of Water

Density of water is a function of temperature, pressure and dissolved ions (conductivity). The density of pure water at 4°C is 1000 kg/m<sup>3</sup>. Groundwater in non-coastal areas has relatively low conductance and does not affect the density of water beyond .001 kg/m<sup>3</sup>. Temperature of the typical range of 5-20°C can affect the density by up to 1 kg/m<sup>3</sup>. The density of water is part of the calculation of groundwater surface elevation as shown in Equation 2. A change of 1 kg/m<sup>3</sup> in density translates to an error of nearly 4mm. This is well below the 1 cm accuracy requirement, therefore the error associated with using the precise density of water is considered negligible. The density of water will be considered constant at 999.0 kg/m<sup>3</sup>.

#### 6.1.1.5 Groundwater Specific Conductivity

The calculation of specific conductivity from actual conductivity and ambient temperature in Equation 7 is a normalization process that compounds the standard errors of both measurements. The uncertainty of individual specific conductivity measurements is calculated as follows:

$$u(SpC_{GW,i}) = \left[ \frac{\partial SpC_{GW,i}^2}{\partial C_{GW,i}} \times u_{A1,C}^2 + \frac{\partial SpC_{GW,i}^2}{\partial T_{GW,i}} \times u_{A1,T}^2 \right]^{\frac{1}{2}} \quad \text{Equation 12}$$

Taking the partial derivatives in Equation 12 become:

$$u(SpC_{GW,i}) = \left[ \left( \frac{1}{1 + 0.0191(T_{GW,i} - 25)} \right)^2 \times u_{A1,C}^2 + \left( \frac{0.0191C_{GW,i}}{(1 + 0.0191(T_{GW,i} - 25))^2} \right)^2 \times u_{A1,T}^2 \right]^{\frac{1}{2}} \quad \text{Equation 13}$$

Where

$u(SpC_{GW,i})$  = uncertainty of individual specific conductivity measurements

### 6.1.1.6 Combined Measurement Uncertainty

#### Conductivity:

Because the only known quantifiable uncertainties are those provided by CVAL, the combined uncertainty of actual conductivity is simply equal to the standard uncertainty values provided by CVAL (See Section 2.4).

#### Specific Conductivity:

Because specific conductivity is derived from the actual conductivity and temperature measurements, which have uncertainties that are provided by CVAL, the uncertainty of the individual specific conductivity measurements is calculated according to Eq. 13 (Section 6.1.1.5).

#### Temperature:

Because the only known quantifiable uncertainties are those provided by CVAL, the combined uncertainty of temperature is simply equal to the standard uncertainty values provided by CVAL (See Section 2.4).

#### Pressure:

Because the only known quantifiable uncertainties are those provided by CVAL, the combined uncertainty of pressure is simply equal to the standard uncertainty values provided by CVAL (See Section 2.4).

#### Groundwater Surface Elevation:

Because groundwater surface elevation is derived from the pressure measurement and the known elevation of the sensor, the combined uncertainty for individual measurements takes into account the positional uncertainty stored in the Named Location Database and the pressure measurement uncertainty provided by CVAL according to Eq. 14.

$$u_c(E_{GW,i}) = \left[ \frac{\partial E_{GW,i}}{\partial E_{sensor}} \times u_c^2(E_{sensor}) + \frac{\partial E_{GW,i}}{\partial P_{GW,i}} \times u_{A1,P}^2 \right]^{\frac{1}{2}} \quad \text{Equation 14}$$

The partial derivatives in Equation 14 become:

$$u_c(E_{GW,i}) = \left[ 1 \times u_c^2(E_{sensor}) + \left( \frac{1000}{\rho g} \right)^2 \times u_{A1,P}^2 \right]^{\frac{1}{2}} \quad \text{Equation 15}$$

Where:

$$u_c(E_{GW,i}) = \text{uncertainty of individual groundwater elevation measurements}$$



## 6.1.2 Uncertainty of the L1 Mean Data Products

The following subsections discuss uncertainties associated with temporally averaged, i.e., L1 mean data products. As stated previously, it is important to note the differences between the *measurement uncertainties* presented in Section 6.1.1 and the uncertainties presented in the following subsections. The uncertainties presented in the following subsections reflect the uncertainty of a time-averaged mean value, that is, they reflect the uncertainty of a distribution of measurements collected under non-controlled conditions (i.e., those found in the field), as well as any uncertainties in the form of *Truth* and *Trueness* related to the accuracy of the field assembly.

### 6.1.2.1 Repeatability (Natural Variation)

To quantify the uncertainty attributable to random effects, the distribution of the individual measurements is used. Specifically, the *estimated standard error of the mean (natural variation)* is computed. This value reflects the repeatability of measurements for a specified time period:

$$u_{NAT}(\bar{X}) = \sqrt{\frac{\sigma^2}{n}} \quad (16)$$

Where

$X$  = measurement, e.g. groundwater pressure, temperature or conductivity

$u_{NAT}(\bar{X})$  = standard error of the mean (natural variation)

$\sigma$  = experimental standard deviation of individual observations for a defined time period

$n$  = number of observations made during the defined period

### 6.1.2.2 Calibration

The calibration uncertainty for a L1 mean DP is similar to that described in Section 6.1.1.1. However, the uncertainties provided by CVAL (see Section 2.4) do not account for i) individual sensor repeatability, or ii) the variation of sensors' responses over a population (reproducibility). These components estimate the uncertainties due to the accuracy of the instrumentation in the form of *Truth* and *Trueness*, a quantity that is not captured by the standard error of the mean. Both values (i.e., conductivity, temperature and pressure) are constant values that will be provided by CVAL and stored in the CI data store. Please refer to AD[10] for further justification regarding evaluation and quantification of this combined uncertainty.

The temperature and conductivity uncertainties provided by CVAL will propagate through to the specific conductivity. This propagation is identical to that shown in 6.1.1.5, however, the uncertainties shown in Eq. (12) and (13) are replaced with  $u_{A3,C}$  and  $u_{A3,T}$  respectively such that:

$$u(\overline{SpC_{GW}}) = \left[ \frac{\partial \overline{SpC_{GW}}^2}{\partial \overline{C_{GW}}} \times u_{A3,C}^2 + \frac{\partial \overline{SpC_{GW}}^2}{\partial \overline{T_{GW}}} \times u_{A3,T}^2 \right]^{\frac{1}{2}} \quad \text{Equation 17}$$

### 6.1.2.3 Combined Uncertainty

The combined uncertainty for L1 AT200 data products are computed by summing the uncertainties from Section 6.1.2.1 and the CVAL provided uncertainties in quadrature:

$$u_c(\overline{X}) = \left( u_{NAT}^2(\overline{X}) + u_{A3,X}^2 \right)^{\frac{1}{2}} \quad (18)$$

Where

$X$  = measurement, e.g. groundwater pressure, temperature or conductivity

**Conductivity:**

$$u_c(\overline{C_{GW}}) = \left[ u_{NAT}^2(\overline{C_{GW}}) + u_{A3,C}^2 \right]^{\frac{1}{2}} \quad \text{Equation 19}$$

**Temperature:**

$$u_c(\overline{T_{GW}}) = \left[ u_{NAT}^2(\overline{T_{GW}}) + u_{A3,T}^2 \right]^{\frac{1}{2}} \quad \text{Equation 20}$$

**Pressure:**

$$u_c(\overline{P_{GW}}) = \left[ u_{NAT}^2(\overline{P_{GW}}) + u_{A3,P}^2 \right]^{\frac{1}{2}} \quad \text{Equation 21}$$

### 6.1.2.4 Groundwater Surface Elevation

The combined uncertainty for groundwater surface elevation includes the combined uncertainties for sensor depth and ground surface. These are discussed above in Section 6.1.1.4.

**Groundwater Surface Elevation:**

$$u_c(\overline{E_{GW}}) = \left[ u_c^2(E_{sensor}) + \left( \frac{1000}{\rho g} \right)^2 u_c^2(\overline{P_{GW}}) \right]^{\frac{1}{2}} \quad \text{Equation 22}$$

### 6.1.2.5 Groundwater Specific Conductivity

The combined uncertainty for specific conductivity includes the uncertainties for temperature and conductivity.

Thus the combined uncertainty for specific conductivity is given as:

$$u_c(\overline{SpC_{GW}}) = \left[ u_{NAT}^2(\overline{SpC_{GW}}) + \left( \frac{\partial \overline{SpC_{GW}}^2}{\partial \overline{C_{GW}}} \right) u_{A3,C}^2 + \left( \frac{\partial \overline{SpC_{GW}}^2}{\partial \overline{T_{GW}}} \right) u_{A3,T}^2 \right]^{\frac{1}{2}} \quad \text{Equation 23}$$

## 6.2 Expanded Uncertainty

We assume the measurement variability is normally distributed (Gaussian). The combined uncertainty represents plus or minus one standard deviation or a confidence interval of 68%. This confidence level is below the industry standard and is therefore expanded to a 95% confidence interval. This is typically calculated by multiplying the combined uncertainty,  $u_{c,r}(y)$ , by a coverage factor,  $k_p$ .

$$U_p = k_p u_{c,r}(y) \tag{Equation 24}$$

Where  $k_{95} = 1.96$  for  $p = 95$  for a given degrees of freedom.  $k_p$  is approximated by a t-distribution with an effective degrees of freedom,  $v_{eff}$ , obtained from the Welch-Satterwaite formula:

$$v_{eff} = \frac{u_c^4(y)}{\sum_{i=1}^n \frac{u_i^4(y)}{v_i}} \tag{Equation 25}$$

The coverage factor then becomes  $k_p = t_p(v_{eff})$ , where  $t_p(v_{eff})$  is obtained from a table.

However to simplify the calculations further, we can conservatively estimate the expanded uncertainty at a 95% confidence level to be two times the combined uncertainty.

$$U_{Expanded} = 2 \times u_c \tag{Equation 26}$$

This expansion is to be applied to all combined uncertainties for the L1 DP described herein.

### 6.3 Uncertainty Budget

The uncertainty budget is a visual aid detailing i) quantifiable sources of uncertainty, ii) means by which they are derived, and iii) the order of their propagation. Uncertainty values denoted in this budget are either derived within this document or are provided by other NEON teams (e.g., CVAL), and stored in the CI data store.

**Table 6-1** Uncertainty budget for individual measurements. Shaded rows denote the order of uncertainty propagation (from lightest to darkest).

Source of Measurement Uncertainty	Measurement uncertainty Component $u(x_i)$	Measurement Uncertainty Value	$\frac{\partial f}{\partial x_i}$	$u_{x_i}(Y) \equiv \left  \frac{\partial f}{\partial x_i} \right  u(x_i)$
Temperature	$u_{A1,T}$	AD[11]	n/a	n/a
Conductivity	$u_{A1,C}$	AD[11]	n/a	n/a
Pressure	$u_{A1,P}$	AD[11]	n/a	n/a

**Table 6-2** Uncertainty budget for L1 mean DPs. Shaded rows denote the order of uncertainty propagation (from lightest to darkest).

Source of Uncertainty	Uncertainty Component $u(x)$	Uncertainty Value	$\frac{\partial f}{\partial x_i}$	$u_{u_{x_i}}(Y) \equiv \left  \frac{\partial f}{\partial x_i} \right  u(x_i)$
Temperature	$u_c(\bar{T}_{GW})$	Eq. (19)	n/a	n/a
Calibration	$u_{A3,T}$	AD[11]	n/a	n/a
Natural variation	$u_{NAT}(\bar{T}_{GW})$		n/a	n/a
Conductivity	$u_c(\bar{C}_{GW})$	Eq. (18)	n/a	n/a
Calibration	$u_{A3,C}$	AD[11]	n/a	n/a
Natural variation	$u_{NAT}(\bar{C}_{GW})$		n/a	n/a
Specific Conductivity	$u_c(\overline{SpC}_{GW})$	Eq. (22)	Eq. 12,13	Eq. 12,13, 22
Temperature	$u_{A3,T}$	AD[11]	n/a	n/a
Conductivity	$u_{A3,C}$	AD[11]	n/a	n/a
Pressure	$u_c(\bar{P}_{GW})$	Eq. (20)	n/a	n/a
Calibration	$u_{A3,P}$	AD[11]	n/a	n/a
Natural variation	$u_{NAT}(\bar{P}_{GW})$		n/a	n/a
Groundwater Surface Elevation	$u_c(\bar{E}_{GW})$	Eq. (21)	n/a	n/a
Pressure	$u_{A3,P}$	AD[11]	n/a	n/a
Sensor Elevation	$u_c(E_{sensor})$	Site metadata	n/a	n/a

## 7 FUTURE PLANS AND MODIFICATIONS

Future system flags may be incorporated into the data stream and included in the QA/QC summary DP ( $Qsum_{5min}$ ) that summarizes any flagged data that went into the computation of the L1 DP.

<i>Title:</i> NEON Algorithm Theoretical Basis Document (ATBD): Groundwater Temperature, Elevation and Specific Conductance		<i>Date:</i> 12/06/2019
<i>NEON Doc. #:</i> NEON.DOC.001328	<i>Author:</i> J. Vance	<i>Revision:</i> B

It is planned that a QA/QC flag for data consistency will be applied according to a developed consistency analysis (AD[05]) and a pass/fail flag will be generated to reflect this activity. Pressure, temperature and conductivity measurements from each well in at a given NEON aquatic site will have the time series data compared against the measurement variance other wells. If a difference between the measurements is less than the defined limits, provided by AQU and maintained in the CI data store, then the sensor will have passed the consistency analysis. Alternatively, a difference outside the defined limits will result in a failed test, and will be flagged as such. L1 DPs that fail the consistency analysis will continue to be reported, but will have an associated failed flag that will be include in the QA/QC summary.

## 8 BIBLIOGRAPHY

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## 9 CHANGELOG